Supplementary Materials for

Seismic detection of strong ground motions by

$M_W5.6$ North Korean nuclear explosion

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North Korean nuclear explosion tests

North Korea conducted 6 underground nuclear explosion (UNE) tests since 2000. The UNE tests were made in 9 October 2006 (1st), 25 May 2009 (2nd), 12 February 2013 (3rd), 6 January 2016 (4th), 9 September 2016 (5th), and 3 September 2017 (6th) (Fig. S1). The seismic magnitudes of the UNEs were $m_b$ 4.3, 4.7, 5.1, 5.1, 5.3, and 6.3 (Hong and Rhie, 2009; Shin et al., 2010; Zhao et al., 2012; Zhang and Wen, 2013; Liu et al., 2018; Tian et al., 2018). The isotropic moment of the first UNE in 2006 was $3.1 \times 10^{14}$ N·m (Koper et al., 2008). The total moment of the second UNE in 2009 was $6.3 \times 10^{15}$ N·m ($M_W$ 4.5) (Ford et al., 2009). The $Lg$ body-wave magnitudes of the five UNEs before the 6th UNE were $m_b(Lg)$ 3.9, 4.5, 4.9, 4.7, and 4.8 (Zhao et al., 2016, 2017). The surface-wave magnitudes of the five UNEs before the 6th UNE were $M_S$ 2.92, 3.65, 3.94, 4.05, and 4.23 (Zhao et al., 2017).

Analysis

The peak ground acceleration (PGA) and peak ground velocity (PGV) attenuation satisfy an equation (Atkinson and Boore, 1995; Marin et al., 2004; Hong et al., 2016):

$$\log G_{i,j,k,l} = A_{i,j,l} + B_{i,j} \log r_{k,l} + C_{i,j} r_{k,l},$$

where $G_{i,j,k,l}$ ($i =$PGA, PGV, $j =$h, v) is the peak ground motion (PGA or PGV) in the horizontal or vertical component at station $k$ for event $l$ at the hypocentral distance of $r_{k,l}$, $A_{i,j,l}$ is a constant calibrated for event size, $B_{i,j}$ is a constant for geometrical spreading, and $C_{i,j}$ is a constant for anelastic absorption. The PGA is in m/s$^2$, the PGV is in m/s, and the distance $r$ is in km.

The constants for the ground motion attenuations were determined using the seismic records for earthquakes in the Korean Peninsula (Hong et al., 2016). The constants for PGA attenuation curves are found to be $B_{PGA,h} = -1.44$, $B_{PGA,v} = -1.54$, $C_{PGA,h} = -0.00211$, and $C_{PGA,v} = -0.00164$, and those for PGV attenuation curves are $B_{PGV,h} = -1.46$, $B_{PGV,v} = -1.52$, $C_{PGV,h} = -0.000939$, and $C_{PGV,v} = -0.000739$. The peak ground motions agree well with the regression curves (Fig. S2). The residuals between the observed PGAs and regression curves are clustered around zero. The source-strength constants satisfy (Hong et
al., 2016)  

\[ A_{PGA,h} = -0.318 + 0.394m_b, \quad A_{PGA,v} = -0.273 + 0.372m_b, \]
\[ A_{PGV,h} = -3.388 + 0.620m_b, \quad A_{PGV,v} = -3.631 + 0.605m_b. \]  

The spatial distribution of ground motions induced by the 2017 UNE is presented in Fig. S3. The PGVs in a hypocentral distance of 1 km are estimated to be 4.46 m/s (0.76-26.01 m/s at a 95 % confidence level) in horizontal components and 2.29 m/s (0.39-13.35 m/s at a 95 % confidence level) in vertical component (Fig. S2). The PGVs in a hypocentral distance of 10 km are estimated to be 0.15 m/s (0.03-0.89 m/s at a 95 % confidence level) in horizontal components and 0.09 m/s (0.02-0.50 m/s at a 95 % confidence level) in vertical component. The PGVs in a hypocentral distance of 100 km are estimated to be 0.004 m/s (0.0007-0.025 m/s at a 95 % confidence level) in horizontal components and 0.003 m/s (0.0005-0.016 m/s at a 95 % confidence level) in vertical component.  

The PGAs in a hypocentral distance of 1 km are estimated to be 169.0 m/s\(^2\) (22.9-1247.8 m/s\(^2\) at a 95 % confidence level) in horizontal components and 150.8 m/s\(^2\) (20.8-1091.7 m/s\(^2\) at a 95 % confidence level) in vertical component. The PGAs in a hypocentral distance of 10 km are estimated to be 5.9 m/s\(^2\) (0.8-43.4 m/s\(^2\) at a 95 % confidence level) in horizontal components and 4.2 m/s\(^2\) (0.6-30.4 m/s\(^2\) at a 95 % confidence level) in vertical component. The PGAs in a hypocentral distance of 100 km are estimated to be 0.14 m/s\(^2\) (0.02-1.02 m/s\(^2\) at a 95 % confidence level) in horizontal components and 0.09 m/s\(^2\) (0.01-0.63 m/s\(^2\) at a 95 % confidence level) in vertical component.  

The dynamic stress changes in a distance of 1 km are 43.5 MPa (7.5-253.9 MPa at a 95 % confidence level) in horizontal component, and 22.3 MPa (3.8-130.3 MPa at a 95 % confidence level) in vertical component. The dynamic stress changes in a distance of 10 km are 1479.7 kPa (253.6-8635.1 kPa at a 95 % confidence level) in horizontal component, and 835.9 kPa (143.2-4878.3 kPa at a 95 % confidence level) in vertical component. The dynamic stress changes in a distance of 100 km are 42.2 kPa (7.2-246.5 kPa at a 95 % confidence level) in horizontal component, and 27.3 kPa (4.7-159.1 kPa at a 95 % confidence level) in vertical component.  

The theoretical ground-motion attenuation curve is compared with the observed ground
motions of the 12 September 2016 $M_L$ 5.8 earthquake in the Korean Peninsula, verifying the validity of the theoretical curve for representation of ground motions in short distances (Fig. S4).

The post-UNE events were relocated using VELHYPO (Fig. S5). The post-UNE events occurred around the flanks of mountains at shallow depths less than 2.5 km (Fig. S6). The 95% confidence ranges of the estimated depth errors vary up to ±1327 m at a 95% confidence level (Fig. S6). The events displayed peak seismic energy around 0.2 Hz (Fig. S7).

We performed long-period waveform inversions of the 23 September 2017 $M_L$ 3.2 post-UNE event. We determined a full moment tensor solution, deviatoric source solution, and double couple solution (Fig. S8). The full moment tensor solution and deviatoric solution are similar, presenting combined sources with substantial strengths of double-couple and compensated linear vector dipole (CLVD) components (Fig. S8). The full moment tensor solution is highly different from the double couple solution. Synthetic waveforms for the full moment tensor solution fit better than those for the double couple solution. The observation suggests that the 23 September 2017 $M_L$ 3.2 post-UNE event may not be a pure natural earthquake. The post-UNE events might occur in weakened medium with possible unidirectional mass displacements (Fig. S9).

References


Hong, T.-K., E. Choi, S. Park, and J. S. Shin (2016), Prediction of ground motion and
dynamic stress change in Baekdusan (Changbaishan) volcano caused by a North Korean nuclear explosion, *Scientific Reports, 6*, 21477, doi:10.1038/srep21477.


Figure S1. Map of six North Korean underground nuclear explosion tests since 2000. The surface topography is presented in color. The figure was created using GMT 4.5.14 (https://www.soest.hawaii.edu/gmt/) and Adobe Illustrator CS6 (http://www.adobe.com/kr/products/illustrator.html).
Figure S2. (a) Horizontal and (b) vertical peak ground accelerations (PGAs) and (c) horizontal and (d) vertical peak ground velocities (PGVs) observed in regional seismic stations for the 6th North Korean underground nuclear explosion in 3 September 2017. The observed PGAs and PGVs agree well with the theoretical curves. The figure was created using Gnuplot 5.0 (https://www.gnuplot.info).
Figure S3. Spatial distribution of (a) horizontal and (b) vertical peak ground accelerations (PGAs) and (c) horizontal and (d) vertical peak ground velocities (PGVs). The locations of stations are marked with triangles. The figure was created using GMT 4.5.14 (https://www.soest.hawaii.edu/gmt/) and Adobe Illustrator CS6 (http://www.adobe.com/kr/products/illustrator.html).
Figure S4. (a) Horizontal and (b) vertical peak ground accelerations (PGAs) and (c) horizontal and (d) vertical peak ground velocities (PGVs) observed in regional seismic stations for the 12 September 2016 $M_L5.8$ earthquake. The observed PGAs and PGVs agree well with the theoretical curves, suggesting the validity of the theoretical curves in short distances. The figure was created using Gnuplot 5.0 (https://www.gnuplot.info).
Figure S5. (a) Map of the 6th North Korean underground nuclear explosion (UNE) and post-explosion events (circles). (b) Comparison between refined and reported locations of post-explosion events. The reported locations were determined by Korea Meteorological Administration (KMA). The event locations were refined using VELHYPO. The figure was created using GMT 4.5.14 (https://www.soest.hawaii.edu/gmt/) and Adobe Illustrator CS6 (http://www.adobe.com.kr/products/illustrator.html).
Figure S6. Vertical source locations of the post-UNE seismic events. The source depths are close to the surface. The figure was created using Gnuplot 5.0 (https://www.gnuplot.info).
Figure S7. Comparison of displacement spectra between the 23 September 2017 M_L 3.2 event and two other post-explosion events in (a) vertical, (b) radial, and (c) tangential components of station YNB. The figure was created using Gnuplot 5.0 (https://www.gnuplot.info).
Figure S8. Long-period waveform inversions of the 23 September 2017 $M_L$3.2 post-explosion seismic event: (a) full moment tensor inversion, (b) deviatoric source inversion, (c) double-couple source inversion, and (d) comparison of inversion results. The full moment tensor solution and deviatoric solution present a combined source behavior with comparable strengths of double-couple and compensated linear vector dipole (CLVD) components. The synthetic waveforms for full moment tensor solution fit better than those for double couple solution. The figure was created using GMT 4.5.14 (https://www.soest.hawai. edu/gmt/) and Adobe Illustrator CS6 (http://www.adobe.com/kr/products/illustrator.html).
Figure S9. An example of shallow source with a considerable CLVD component. (a) A simplified tunnel model with the horizontal and vertical lengths of tunnel entrance to be $\xi_H$ and $\xi_V$. The collapsed-tunnel length is $\xi_L$, the thickness of collapsed medium is $r$, and the depth of tunnel is $d$. (b) A schematic model of vertical tunnel collapse. The figure was created using Adobe Illustrator CS6 (http://www.adobe.com/kr/products/illustrator.html).