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# A WAVELET-BASED METHOD FOR SIMULATION OF SEISMIC WAVE PROPAGATION

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Seismic wave propagation (e.g., both *P-SV* and *SH* in 2-D) can be modeled using wavelets. The governing elastic wave equations are transformed to a first-order differential equation system in time with a displacement-velocity formulation. Spatial derivatives are represented with a wavelet expansion using a semigroup approach. The evolution equations in time are derived from a Taylor expansion in terms of wavelet operators. The wavelet representation allows high accuracy for the spatial derivatives. Absorbing boundary conditions are implemented by including attenuation terms in the formulation of the equations. The tractionfree condition at a free surface can be introduced with an equivalent force system. Irregular boundaries can be handled through a remapping of the coordinate system. The method is based on a displacement-velocity scheme which reduces memory requirements by about 30% compared to the use of velocity-stress. The new approach gives excellent agreement with analytic results for simple models including the Rayleigh waves at a free surface. A major strength of the wavelet approach is that the formulation can be employed for highly heterogeneous media and so can be used for complex situations.



### Wavelets & Applications (I-1)

### Signal processing

- Phase picking (P, S, Lg)
- Estimation of a time-varying spectral density matrix for three component seismic data
- Measurement of the anisotrophy at a given area using a relative phase between components

#### Numerical analysis

- Simplification of governing equations by applying a wavelet function
- Composite use of a FD scheme and an interpolating wavelet transform
- Application of wavelets in differentiation of functions

### Wavelets & Applications (I-2)





### Wavelets & Applications (I-3)



#### Discrete wavelet transform (DWT)

- Daubechies (1992) built the foundation of a DWT
- Based on a multiresolution analysis that decomposes a signal into components of different scales

#### Wavelets & parabolic PDEs

- Beylkin & Keiser (1997) introduced a numerical scheme for parabolic PDEs using semigroup approach and wavelets
- Semigroup approach

$$\partial_t g = \mathcal{L}g + \mathcal{N}f(g)$$
  
$$\Rightarrow g_{n+1} = e^{\delta t \mathcal{L}} g_n + \delta t \left(\gamma N_{n+1} + \sum_{m=0}^{R-1} \beta_m N_{n-m}\right)$$



- Spatial derivative operators in elastic wave equations are treated through wavelet transforms in a physical domain
- The resulting second order differential equations for time evolution are solved via a system of first-order differential equations using a displacement-velocity formulation, which can reduce memory requirements by 30 % compared to a velocity-stress formulation
- With the combined aid of a semigroup representation and spatial differentiation using wavelets, a uniform numerical accuracy of spatial differentiation can be maintained across the domain

### A Wavelet-Based Method (II-2)



### Numerical difficulties

- Numerical formulation of wave equations for the parabolic PDE scheme
- Representation of matrix operators in wavelet bases
- Treatment of artificial boundary conditions
- Treatment of inherent periodic boundary conditions
- Implementation of external boundary conditions (e.g. free-surface conditions, rigid boundary conditions)
- Expansion to complex structure problems

### A Wavelet-Based Method (II-3)



#### Corresponding remedies

- Introduction of a displacement-velocity formulation
- Application of a Taylor expansion for the simplification of matrix operators
- Introduction of absorbing boundary conditions
- Adjustment of physical parameters around the artificial boundaries
- Application of equivalent force terms in the system of equations for external boundary conditions
- Introduction of grid generation scheme



### **Formulation of Equations (III)**

#### ■ 2-D elastic wave equations (*P-SV*)

$$\frac{\partial^{2} u_{x}}{\partial t^{2}} = \frac{1}{\rho} \left( \frac{\partial \sigma_{xx}}{\partial x} + \frac{\partial \sigma_{xz}}{\partial z} + f_{x} \right) \\ \frac{\partial^{2} u_{z}}{\partial t^{2}} = \frac{1}{\rho} \left( \frac{\partial \sigma_{xz}}{\partial x} + \frac{\partial \sigma_{zxz}}{\partial z} + f_{z} \right) \right\} \Rightarrow \partial_{t} \mathbf{U} = \mathbf{L} \mathbf{U} + \mathbf{F}$$

where,

$$\mathbf{U} = \begin{pmatrix} u_{x} \\ v_{x} \\ u_{z} \\ v_{z} \end{pmatrix} \quad \mathbf{L} = \begin{pmatrix} 0 & I & 0 & 0 \\ L_{xx} & 0 & L_{xz} & 0 \\ 0 & 0 & 0 & I \\ L_{zx} & 0 & L_{zz} & 0 \end{pmatrix} \quad \mathbf{F} = \frac{1}{\rho} \begin{pmatrix} 0 \\ f_{x} \\ 0 \\ f_{z} \end{pmatrix}$$

- **U** : vector variable
- L : matrix operator
- **F** : force vector

### **Test of Equivalent Forces (IV)**







The equivalent force terms for boundary conditions can simulate exactly the effects which are caused by the presence of the boundaries



### **Unbounded Medium (V)**





### Media with a Free Surface (VI-1)



SH waves



### Media with a Free Surface (VI-2)





### Media with a Free Surface (VI-3)



#### Snapshots of wavefields



### **Two-Layered Media (VII-1)**









#### Snapshots of SH wavefields







#### Snapshots of *P-SV* wavefields



### **Gradient Velocity Media (VIII-1)**



#### Model & Time responses



β: 1.8~4.5 km/s

surface

### Gradient Velocity Media (VIII-2)





# Media with Topographic Surface (IX-1)

Application to topographic problems

 Introduce a grid generation scheme : mapping a rectangular grid system to a curved grid system with consideration of physical topography

#### Numerical model $S_3$ $S_2$ A homogeneous medium with a sinusoidal free 2500 z (m) $\alpha$ = 3.5 km/s $\beta = 2.0$ km/s surface $\rho = 2.2 \text{ g/cm}^3$ Three different positions 5000 of sources; at a trough, a hill and a crest beneath a 0 2500 5000 7500 10000 free surface x (m)



#### Snapshots of *P-SV* wavefields (t=1.9 s)







#### Numerical models



Pointwise random heterogeneous medium with Gaussian probability distribution and standard deviation of 20 % in velocity

Stochastic heterogeneous medium using a Von Karman autocorrelation function with a correlation distance of 5 km



#### Snapshots of *P-SV* wavefields (t=1.9 s, X comp.)



x (m)

x (m)

x (m)



#### **Time responses (one-layered media, X comp.)**

Homogeneous

• Random heterog.

#### Von Karman ACF



# Media with Cavities (XI-1)



#### Numerical model



#### One or five cavities inside a medium

 Implementation of boundary conditions

-via additional equivalent force terms (i.e. traction-free boundary conditions)
-can implement various boundary conditions at various locations at the same time





#### Snapshots in one cavity model (Z comp.)





### Media with Cavities (XI-3)

#### **Snapshots in five cavities model (***Z* **comp.)**



### **Tool for Quantitative Study (XII-1)**



Comparison of accuracy with 4<sup>th</sup>-order FDM
 Apply the first-order differentiation to the input signal which is corresponding to the quick variation of physical parameters in random heterogeneous media

15 40 analytic input signal analytic Ж 10 4th-order FD wavelet  $\overline{}$ 5 20 0 0 -5 -10 -20 -15 5 15 0 10 20 -40 where,  $N_x = 64$ 

0

10

5

15

20

# **Tool for Quantitative Study (XII-2)**



### Application to Q estimation



#### **Cf. Other Researches**

-Frankel & Clayton (1986) :  $\theta$ min = 30-45°

-Roth & Korn (1993) :  $\theta_{\min} = 20-40^{\circ}$ 

-Jannaud et al. (1991) :  $\theta_{\min} = 90^{\circ}$ 



- Can obtain high accuracy of numerical responses of media for wave propagation due to almost no loss of accuracy during spatial differentiation and third order accuracy for time derivatives
- Stable in highly perturbed velocity media Numerically stable scheme
- Can implement various complex boundary conditions at various locations easily using equivalent force terms
- Energy is conserved 

   can be applied in seismic quantitative studies (e.g. estimation of amount of energy loss during wave propagation)